



# SYE 3801 Aerodynamics Spring 2015

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# Hydrostatic Equation: Review

- Absolute Altitude ( $h_a$ ) is important especially for space flight, because the local acceleration due to gravity at *sea level*:

$$g = g_0[r/h_a]^2 = g_0[r/(r + h_G)]^2 \quad (1) \leftarrow \text{Hydrostatic Equation}$$

Consider a stationary fluid element:

Volume of element:  $1 \times 1 \times dh_G$

Mass of element:  $\rho \times 1 \times 1 \times dh_G$

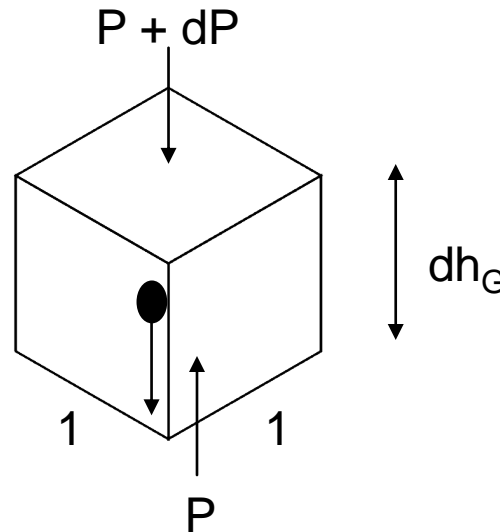
Weight of element:  $g \times \rho \times dh_G$

$$F = P \times A = P \times 1 \times 1$$

Three forces are balanced:

$$P = P + dP + \rho g dh_G$$

$$dP = -\rho g dh_G \quad (2) \leftarrow \text{Hydrostatic Equation (valid for gases and liquids)}$$



# Geopotential Altitude

- The value of  $g$  changes with altitude. In calculations, a constant value of  $g_0$  is used ( $g_0 = g$  at sea level). To compensate for changing  $g$  with altitude, we use geopotential altitude.

- Geopotential Altitude ( $h$ ) is given by:  $h = \{r/(r + h_G)\}h_G$  (3)

- $h \approx h_G$  for low altitudes ( $< 65\text{km}$ ,  $213,000$  ft)

- Divide the Hydrostatic Equation by the Equation of State:

- $$\frac{dP}{P} = -\frac{\rho g_0 dh}{\rho RT} = -\frac{g_0}{RT} dh$$

- We can obtain pressure  $P$  at  $h$  by integrating between  $h_1$  and  $h$ :

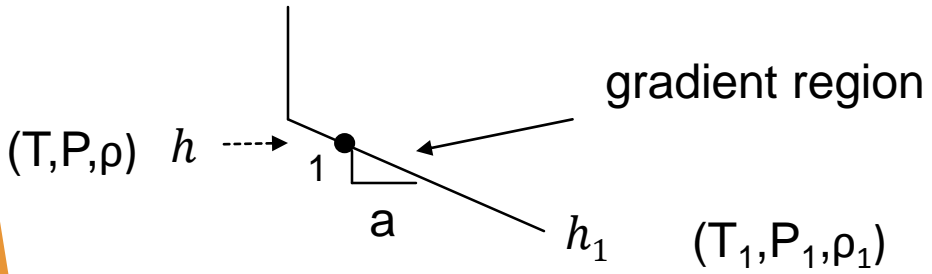
$$\int_{P_1}^P \frac{dP}{P} = -\frac{g_0}{RT} \int_{h_1}^h dh$$
$$\ln\left(\frac{P}{P_1}\right) = -\frac{g_0}{RT} (h - h_1)$$

$$\frac{P}{P_1} = e^{-\frac{g_0}{RT}(h-h_1)}, \text{for isothermal region} \quad (4)$$

- From the Equation of State:

$$\frac{P}{P_1} = \frac{\rho T}{\rho_1 T_1} = \frac{\rho}{\rho_1}$$

$$\frac{\rho}{\rho_1} = e^{-\frac{g_0}{RT}(h-h_1)}, \text{ for isothermal region (5)}$$



$$\frac{T - T_1}{h - h_1} = \frac{dT}{dh} = a$$

- We know that

$$\frac{dP}{P} = -\frac{\rho g_0 dh}{\rho RT} = -\frac{g_0}{RT} dh = -\frac{g_0}{aR} \frac{dT}{T}$$

$$\int_{P_1}^P \frac{dP}{P} = -\frac{g_0}{aR} \int_{T_1}^T \frac{dT}{T}$$

$$\ln\left(\frac{P}{P_1}\right) = -\frac{g_0}{aR} \ln\left(\frac{T}{T_1}\right)$$

Recall:  $\ln(A) - \ln(B) = \ln(A/B)$ , and  $q \ln(A) = \ln(A^q)$

$$\ln\left(\frac{P}{P_1}\right) = -\frac{g_0}{aR} \ln\left(\frac{T}{T_1}\right) \rightarrow \ln\left(\frac{P}{P_1} / \frac{T}{T_1}^{-\frac{g_0}{aR}}\right) = 0$$

Recall:  $\ln(A) = B \rightarrow e^B = A$

$$e^0 = \frac{P}{P_1} / \frac{T}{T_1}^{-\frac{g_0}{aR}} = 1 \rightarrow \frac{P}{P_1} = \left[\frac{T}{T_1}\right]^{-\frac{g_0}{aR}}$$

**(6)** For gradient region

- From the Equation of State:

$$\frac{P}{P_1} = \frac{\rho RT}{\rho_1 R T_1} = \left[ \frac{T}{T_1} \right]^{-\frac{g_0}{aR}}$$

$$\frac{\rho}{\rho_1} = \left[ \frac{T}{T_1} \right]^{-\left[ \frac{g_0}{aR} + 1 \right]} \quad (7) \text{ For gradient region}$$

- Variation of T is linear with altitude:

$$T = T_1 + a(h - h_1) \quad (8)$$

Recall: P,  $\rho$ , and T under standard conditions (sea level):

$$P_s = 1.01325 \times 10^5 \text{ N/m}^2 = 1 \text{ atm}$$

$$\rho_s = 1.225 \text{ kg/m}^3$$

$$T_s = 288.15 \text{ K}$$

- Recall:  $P$ ,  $\rho$ , and  $T$  under standard conditions (sea level):

$$P_s = 1.01325 \times 10^5 \text{ N/m}^2 = 1 \text{ atm}$$

$$\rho_s = 1.225 \text{ kg/m}^3$$

$$T_s = 288.15 \text{ K}$$

- Using equations 1 – 8 the standard atmosphere tables are generated.
- Note: the actual atmosphere data ( $T, P, \rho$ ) at any given altitude may be different from the data given in tables depending on the day.

Example: Calculate the standard atmosphere values of  $T, P$ , and  $\rho$  at A geopotential altitude of 14 km.

Beginning at sea-level, the first region is a gradient region from  $h = 0$  to  $h = 11$  km.

$$a = \frac{dT}{dh} = \frac{216.66 - 288.15}{11 - 0} = -6.5 \frac{K}{km} = -0.0065 \frac{K}{m}$$

At sea-level  $P_1 = 1.01 \times 10^5 \text{ N/m}^2$ ,  $\rho_1 = 1.23 \text{ kg/m}^3$ ,  $g_0 = 9.8 \text{ m/s}^2$   
 $R = 287 \text{ J/(KgK)}$

Using equation (6) for the gradient region:

$$P = P_1 \times \left[ \frac{T}{T_1} \right]^{-\frac{g_0}{aR}} = (1.01 \times 10^5) \left[ \frac{216.66}{288.15} \right]^{\frac{-9.8}{-0.0065 \times 287}}$$

$$P(\text{at } 11 \text{ km}) = 2.26 \times 10^4 \frac{N}{m^2}$$

Using equation (7) for the gradient region:

$$\rho = \rho_1 \times \left[ \frac{T}{T_1} \right]^{-\left\{ \frac{g_0}{aR} + 1 \right\}} = (1.23) \left[ \frac{216.66}{288.15} \right]^{-\left[ \frac{9.8}{-0.0065 \times 287} + 1 \right]}$$

$$= 0.367 \frac{kg}{m^3} (\text{at } 11 \text{ km})$$



- The values of  $P$  and  $\rho$  form the base for the first isothermal region. For the isothermal region use equations (4) and (5):

$$P = P_1 e^{-\frac{g_0(h-h_1)}{RT}} = (2.26 \times 10^4) e^{-\left[\frac{9.8}{(287)(216.66)}\right](14,000-11,000)} = 1.41 \times 10^4 \frac{N}{m^2}$$

$$\rho = \rho_1 * \frac{P}{P_1} = 0.367 * \frac{1.41 \times 10^4}{2.26 \times 10^4} = 0.23 \frac{kg}{m^3}$$